

Improving Grounding Line Discretization using an Embedded-Boundary Approach in BISICLES

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U.S. DEPARTMENT OF
ENERGY

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Science

BISICLES

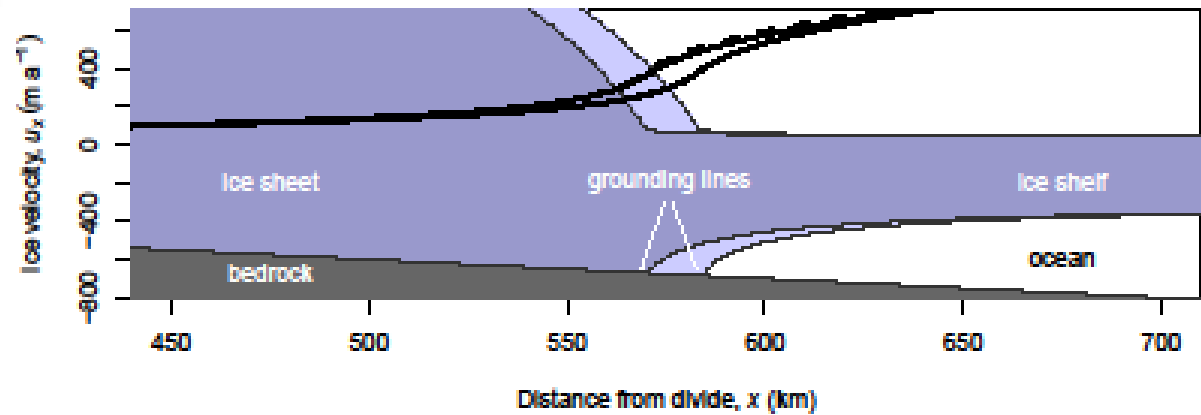
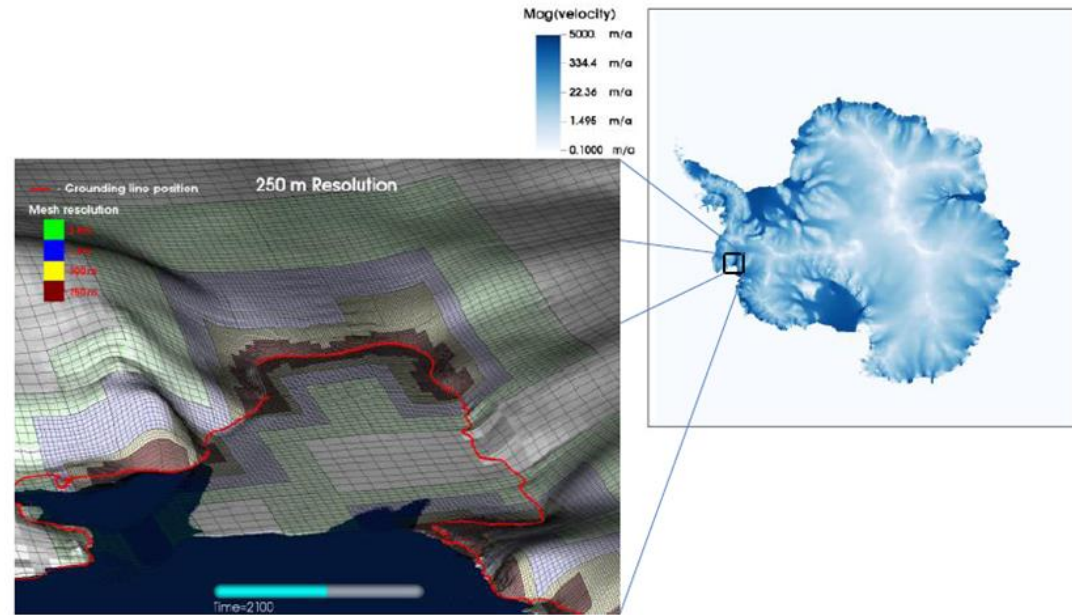
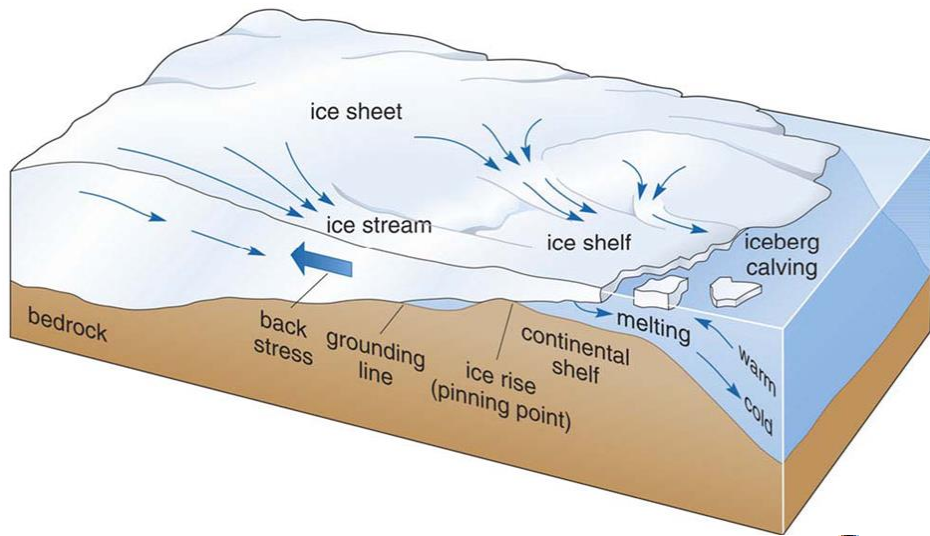


Joint work with:

- ❑ **Peter Schwartz (LBNL)**
- ❑ **Phil Colella (LBNL)**
- ❑ **Stephen Cornford (Bristol)**
- ❑ **Mark Adams (LBNL)**
- ❑ **Esmond Ng (LBNL)**



Land Ice Sheets - coupling with Oceans



Motivation: Projecting future Sea Level Rise

- ❑ Potentially large Antarctic contributions to SLR resulting from marine ice sheet instability, particularly from WAIS.
- ❑ Climate driver: subshelf melting driven by warm(ing) ocean water intruding into subshelf cavities.
- ❑ Melt-driven thinning, loss of shelf buttressing lead to grounding-line retreat.
- ❑ Paleorecord implies that WAIS has deglaciated in the past.

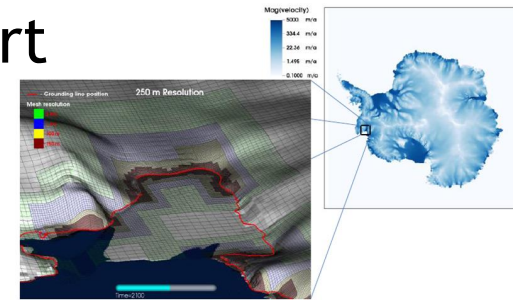


DOE Context - PISCEES and ACME

Part of the DOE “big picture” in climate

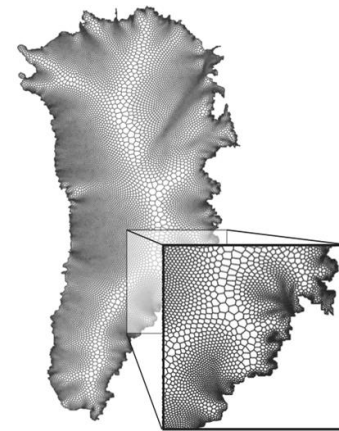
❑ PISCEES (Predicting Ice Sheet and Climate Evolution at Extreme Scales)

- DOE-sponsored (SciDAC2) ice-sheet modeling effort
- Leverages DOE modeling, HPC capabilities
- Dycore development
 - BISICLES - block-structured finite-volume AMR, L1L2
 - FELIX - Finite Element unstructured mesh, Blatter-Pattyn/Stokes
- Initialization, UQ, V&V



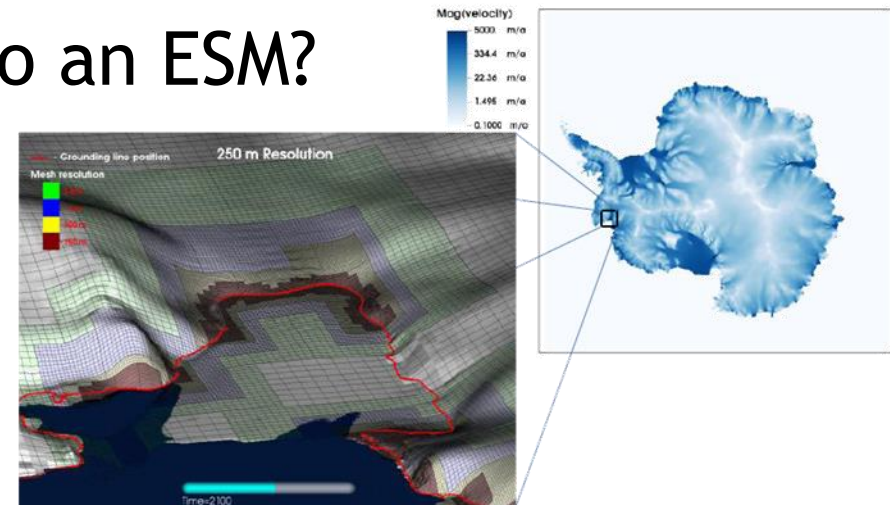
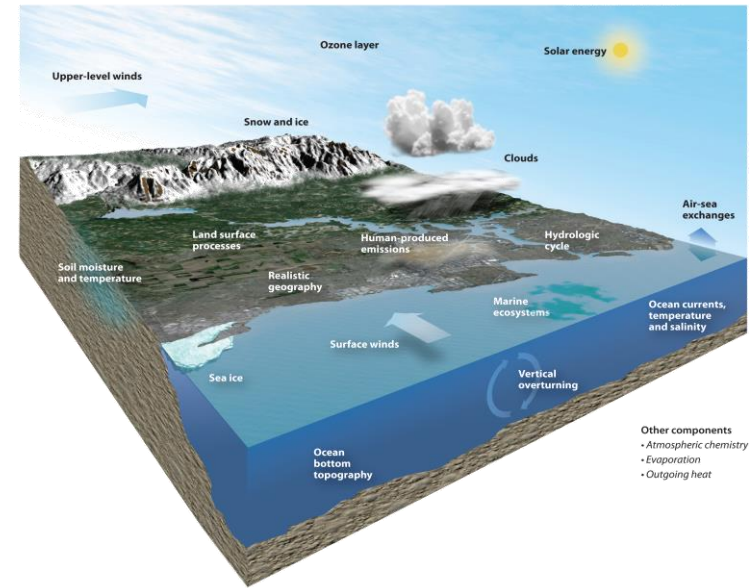
❑ ACME (Accelerated Climate Model for Energy)

- DOE-sponsored ESM effort
 - 3 science questions (#3 is cryospheric contribution to SLR)
- Starting point is CESM



Big Picture -- target

- Aiming for coupled ice-sheet-ocean modeling in ESM
- Multi-decadal to century timescales
- Target resolution:
 - Ocean: 0.1 Degree
 - Ice-sheet: 500 m (adaptive)
- Why put an ice-sheet model into an ESM?
 - fuller picture of sea-level change
 - feedbacks may matter on timescales of years, not millenia
- Credible projections require correct GL dynamics



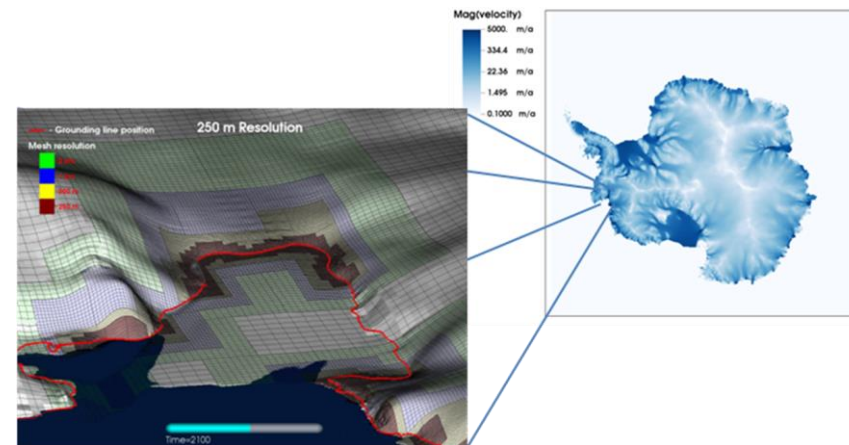
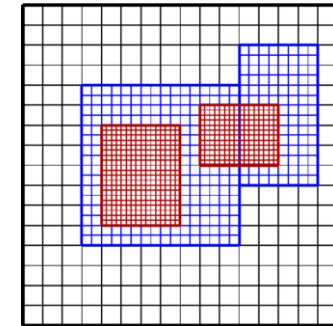
Grounding-line dynamics experiments

- ❑ Series of ice-sheet modeling community model intercomparison projects designed to understand issues in modeling of GLs
 - MISMIP, MISMIP3D, MISMIP+
- ❑ All point to a need for very fine spatial resolution to get GL dynamics right (sub-km in most cases)
- ❑ Prime use case for adaptive mesh refinement (AMR)



BISICLES Ice Sheet Model

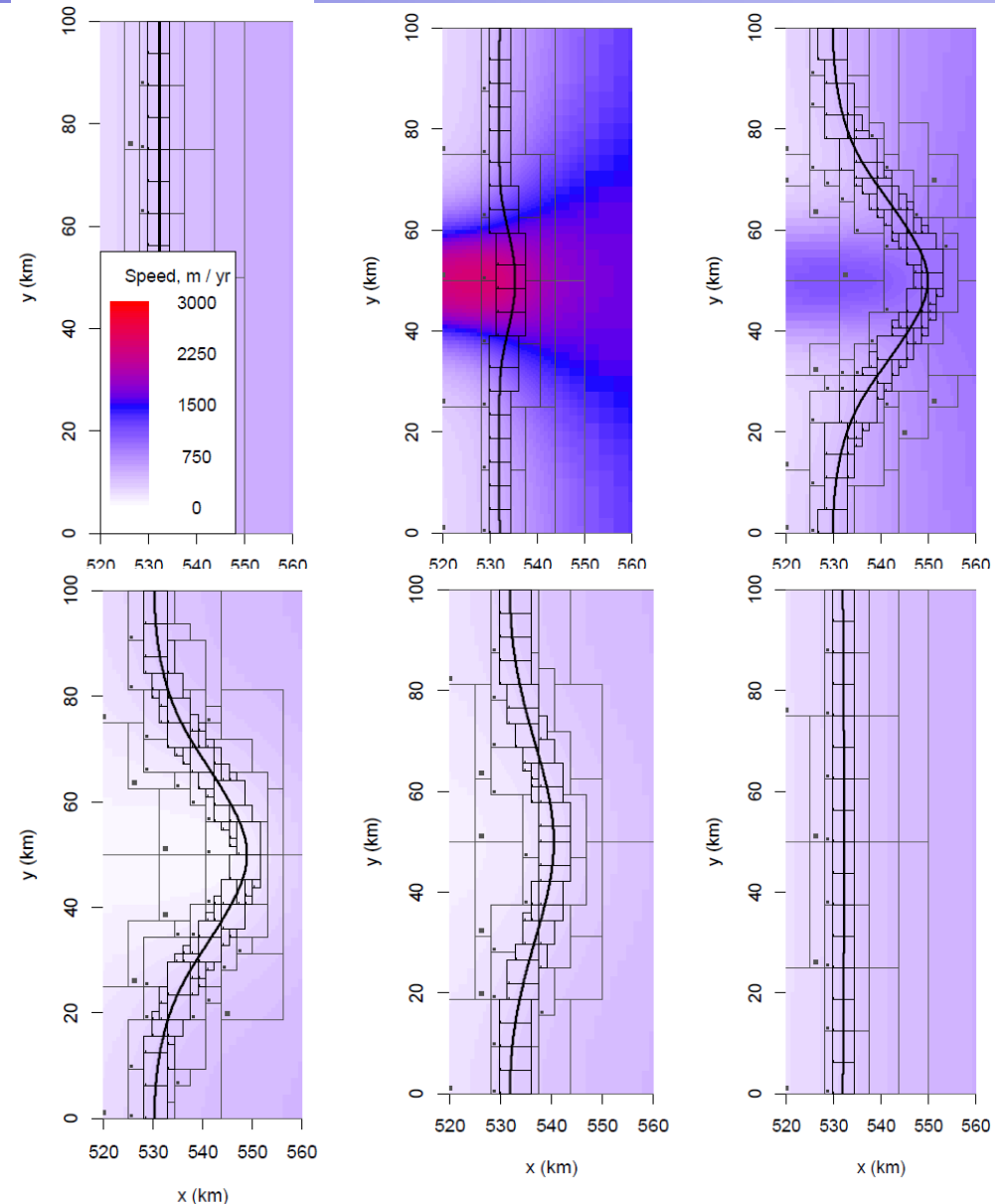
- ❑ Scalable adaptive mesh refinement (AMR) ice sheet model
 - Dynamic local refinement of mesh to improve accuracy
- ❑ Chombo AMR framework for block-structured AMR
 - Support for AMR discretizations
 - Scalable solvers
 - Developed at LBNL
 - DOE ASCR supported (FASTMath)
- ❑ Collaboration with Bristol (U.K.) and LANL
- ❑ Variant of “L1L2” model (Schoof and Hindmarsh, 2009)
- ❑ Coupled to Community Ice Sheet Model (CISM).
- ❑ Users in Berkeley, Bristol, Beijing, Brussels, and Berlin...



BISICLES Results - MISMIP3D

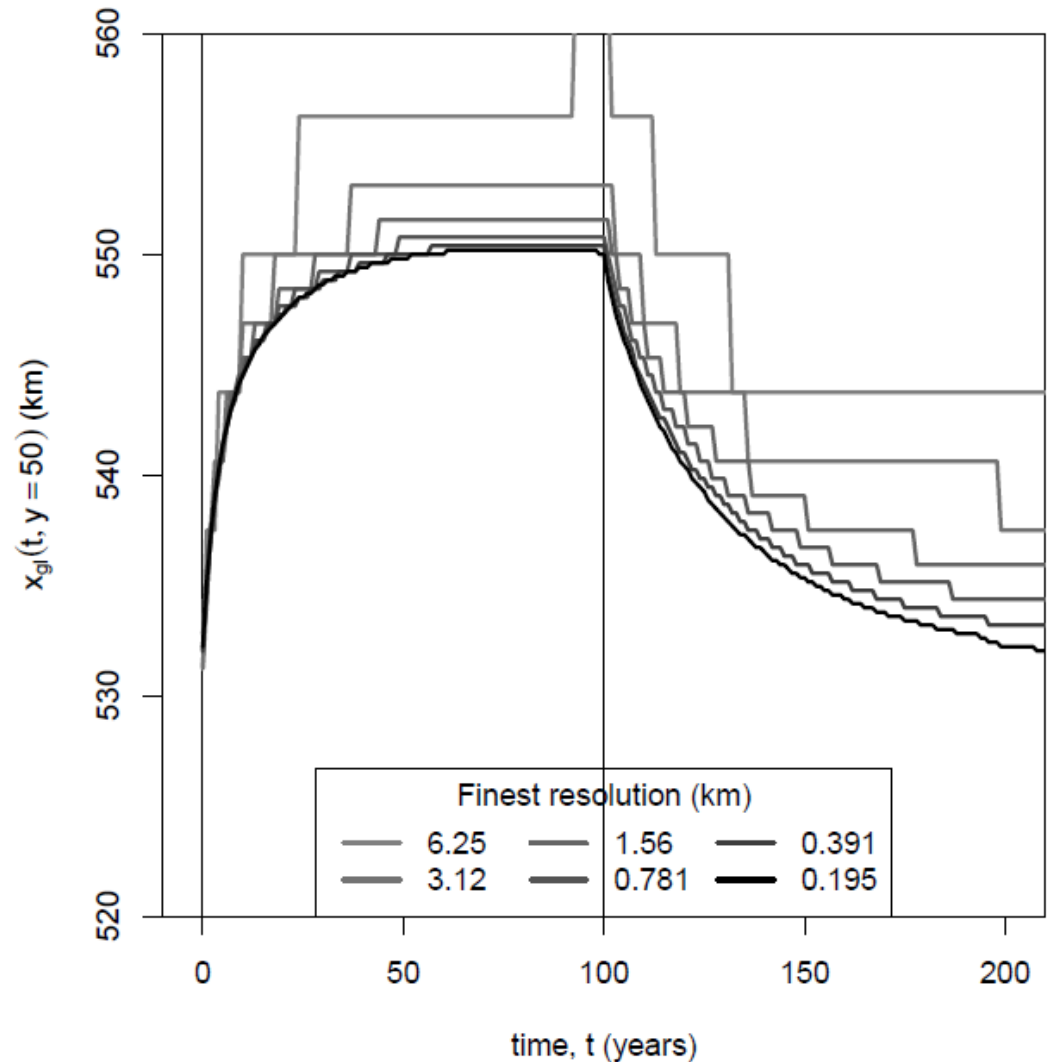
Experiment P75R: (Pattyn et al (2011))

- ❑ Begin with steady-state (equilibrium) grounding line.
- ❑ Add Gaussian slippery spot perturbation at center of grounding line
- ❑ Ice velocity increases, GL advances.
- ❑ After 100 years, remove perturbation.
- ❑ Grounding line should return to original steady state.
- ❑ Figures show AMR calculation:
 - $\Delta x_0 = 6.5km$ base mesh,
 - 5 levels of refinement
 - Finest mesh $\Delta x_4 = 0.195km$.
 - $t = 0, 1, 50, 101, 120, 200 yr$
- ❑ Boxes show patches of refined mesh.
- ❑ GL positions match Elmer (full-Stokes)

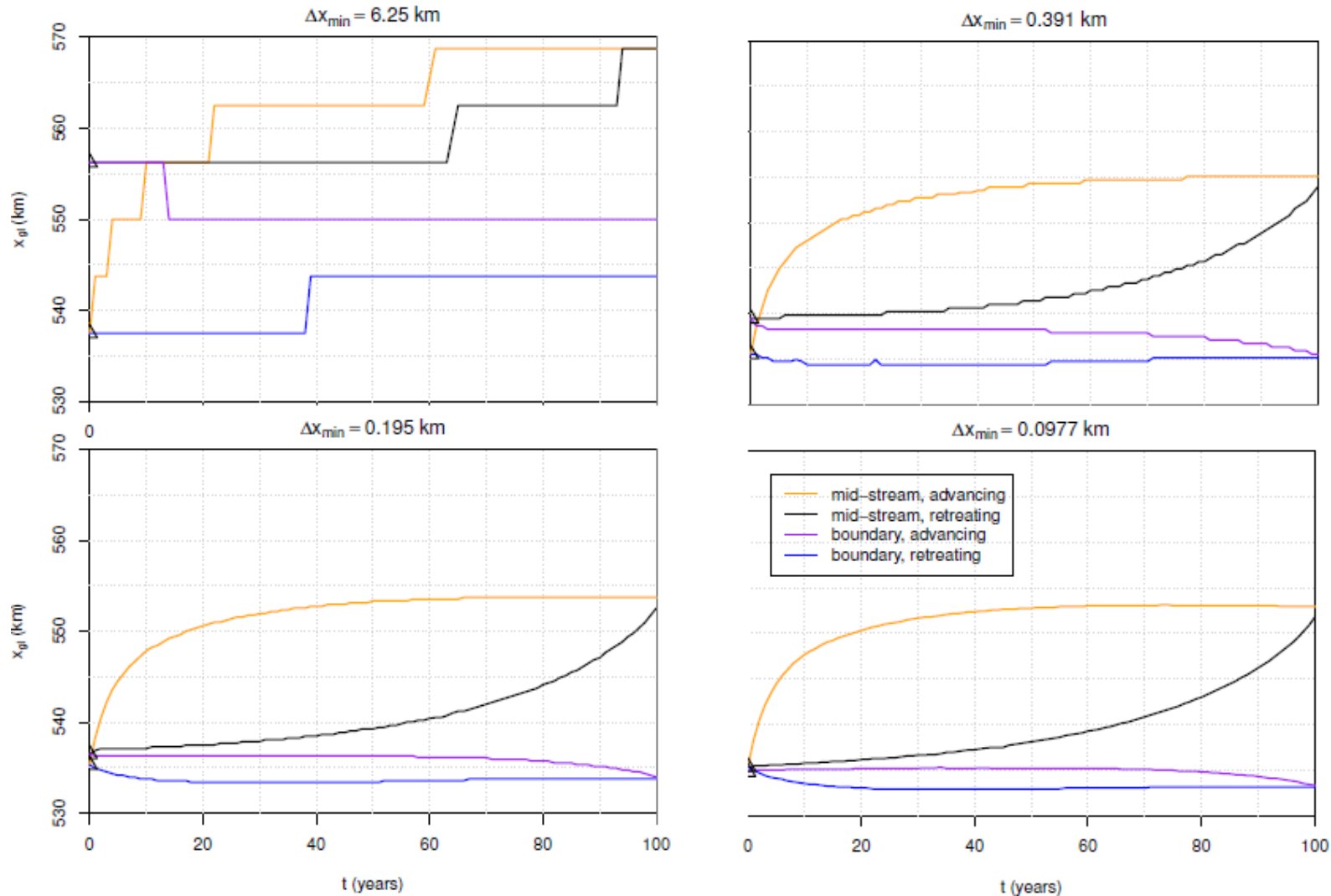


MISMIP3D: Mesh resolution

- Plot shows grounding line position x_{GL} at $y = 50\text{km}$ vs. time for different spatial resolutions.
- $\Delta x = 0.195\text{km} \rightarrow 6.25\text{ km}$
- Appears to require finer than 1 km mesh to resolve dynamics
- Converges as $O(\Delta x)$ (as expected)



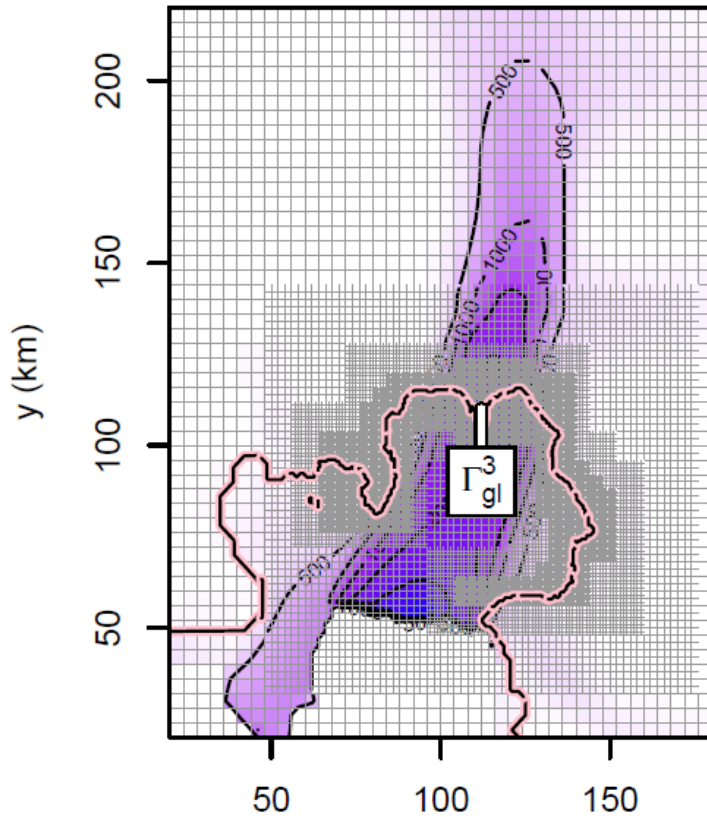
MISMIP3D (cont): Spatial Resolution



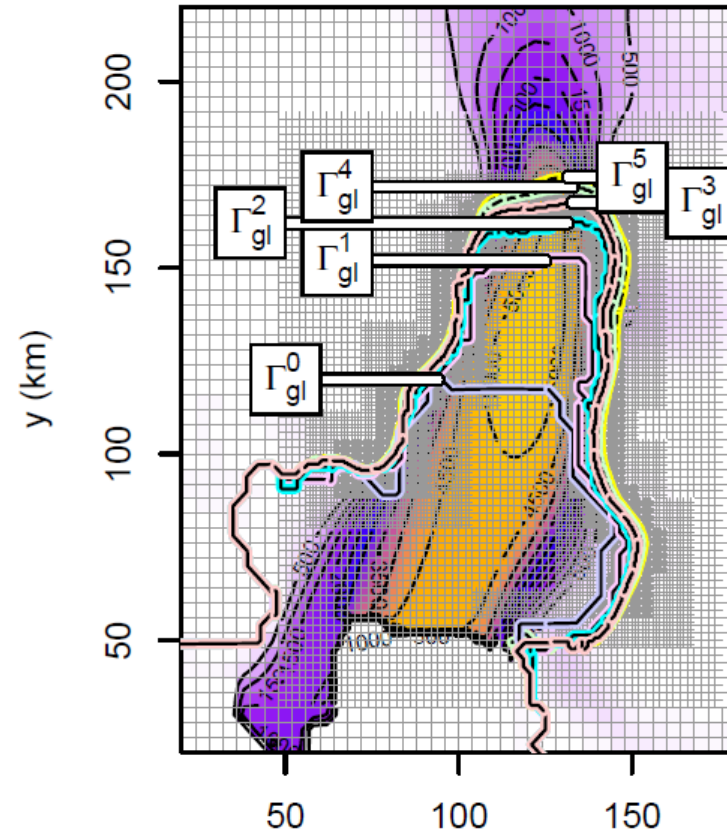
- Very fine (~200 m) resolution needed to achieve full reversibility!



Pine Island Glacier (Cornford, Martin, et al, JCP)



Initial Condition



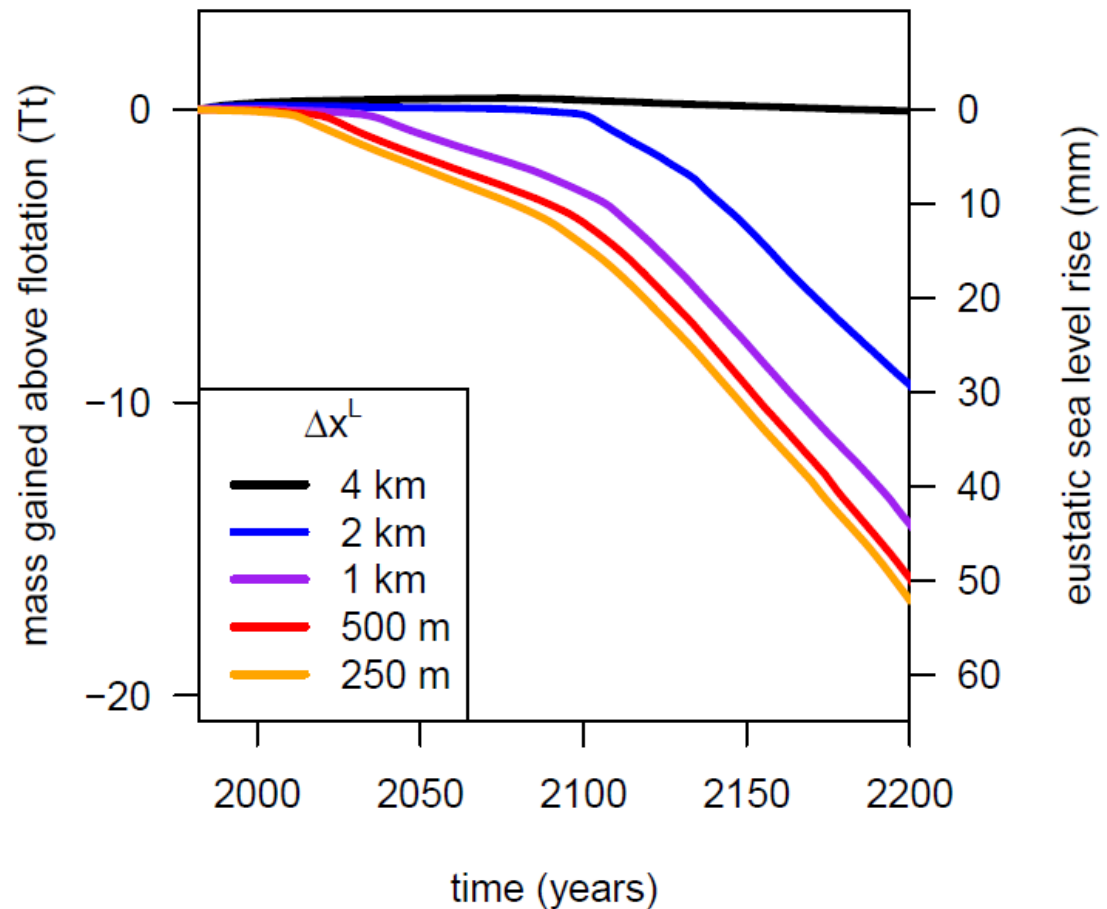
Solution after 30 years

Coloring is ice velocity, Γ_{gl} is the grounding line. Superscripts denote number of refinements. Note resolution-dependence of Γ_{gl}

Amundsen Sea (Cornford, Martin et al, submitted)

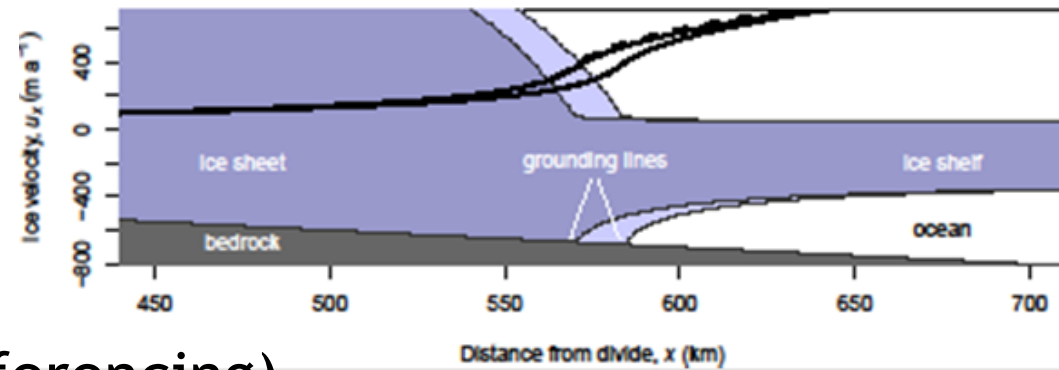
- Need at least 2 km resolution to get any measurable contribution to SLR.
- Sub-km is better.
- Appears to converge at first-order in Δx

SLR vs. year, Amundsen Sea Sector



GL Resolution requirements

- ❑ Not model-specific; reported by many authors
 - Full-Stokes (Elmer - Durand et al)
 - Hybrid SSA-SIA (PISM-PIK)
- ❑ Such resolution requirements are inconvenient, at best.
- ❑ Point to the fact that in models with hydrostatic formulations, GL is a singular point (set)
 - Basal friction drops to zero
 - SSA-type equations go from parabolic to elliptic
 - Surface slopes are discontinuous (one-sided differencing)



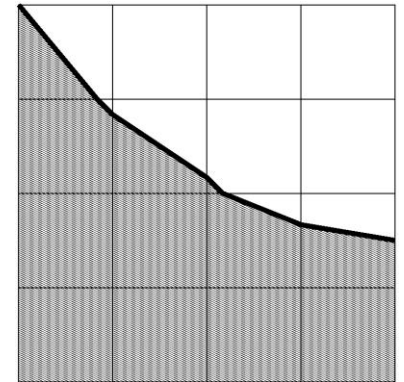
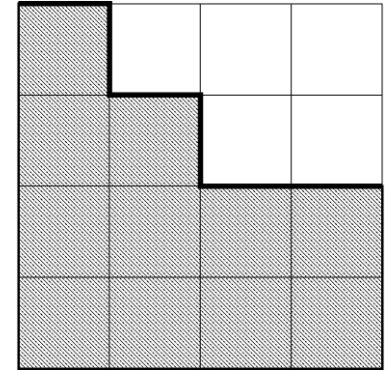
Other approaches

- ❑ Sub-km mesh resolution requirements are inconvenient at best for continental-scale models.
- ❑ Many attempts to handle this through subgrid-scale models
 - Transition zones (Pattyn)
 - Partial-cell parameterization (Gladstone et al, Seroussi et al)
 - More complicated asymptotics - (Leguy et al)



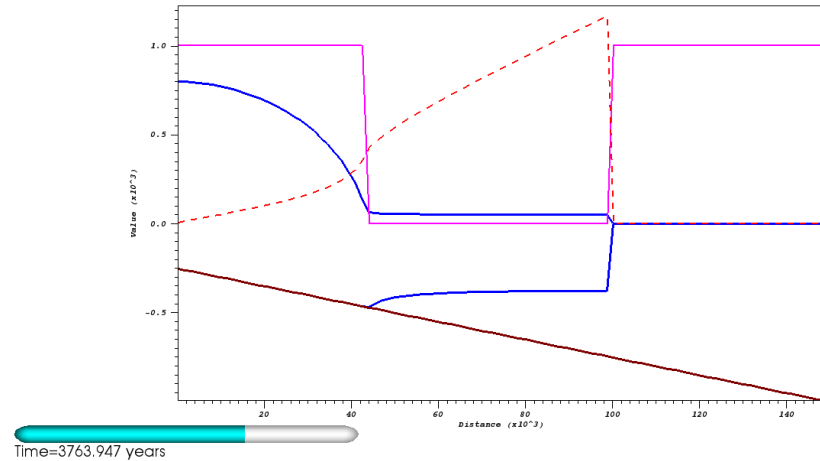
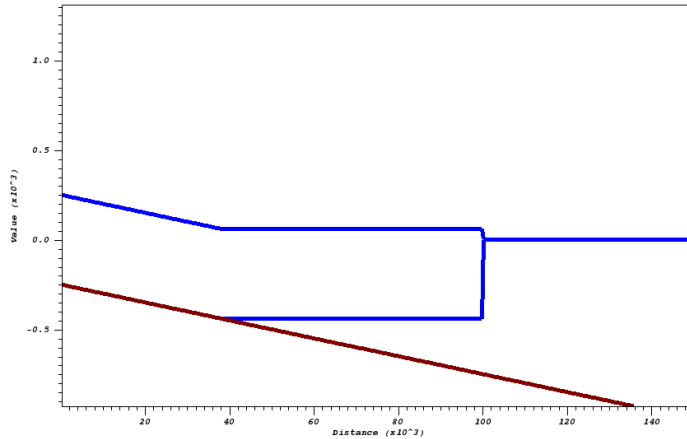
Embedded Boundary (EB) for Grounding Lines

- **Embedded Boundary (EBChombo)**
 - Currently force GL and ice margins to cell faces
 - “Stair-step” discretization
Known to be inadequate from experience with Stefan Problem in other contexts!
 - Use Chombo Embedded-boundary support to improve discretization of GL’s and ice margins.
 - Can solve as a Stefan Problem, with appropriate jump conditions enforced at grounding line.
(as in Schoof, 2007)



Flowline (1D) model problem

- Based on Vieli and Payne (2005)



- SSA Momentum balance reduces to:

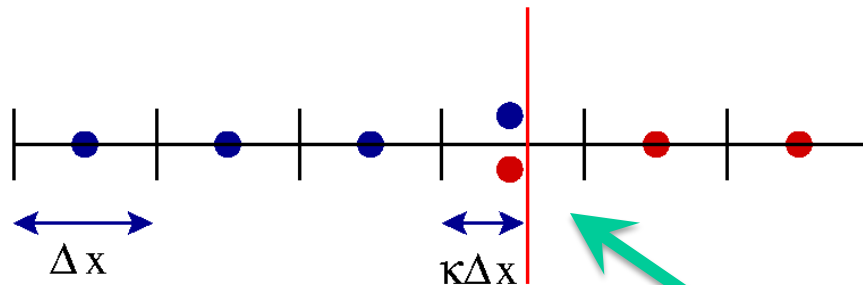
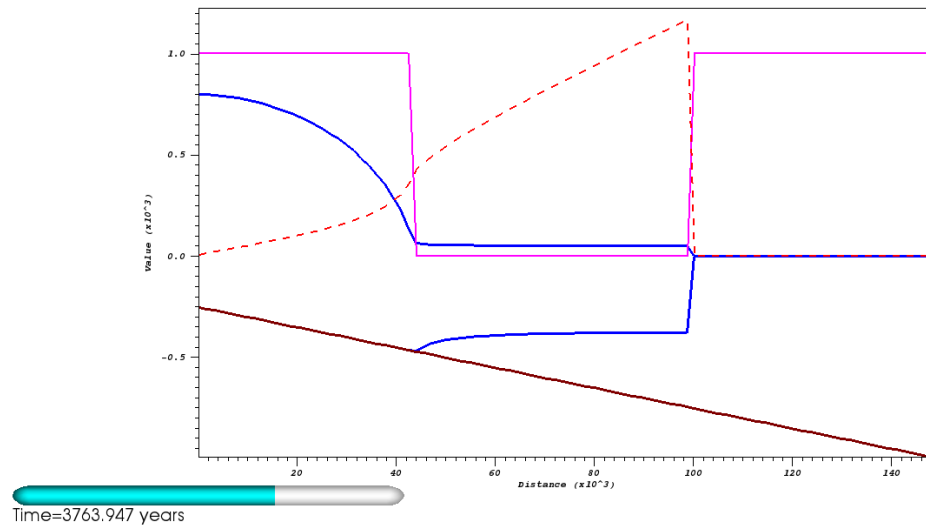
$$\left(\beta - \frac{\partial}{\partial x} \left(4\mu H \frac{\partial}{\partial x} \right) \right) u_b = -\rho g H \frac{\partial s}{\partial x}$$

- Mass Conservation reduces to:

$$\frac{\partial H}{\partial t} + \nabla \cdot (uH) = Src$$

Multifluid formulation

- ❑ Can conceive of the grounding line problem as a phase-change across a multifluid interface (Stefan problem)
- ❑ Discretization follows Crockett, Colella, and Graves (2011)



Multivalued cell

Multifluid Velocity Solve

- ❑ Multifluid discretization (Crockett et al, 2011)
- ❑ Grounded, floating “phases” discretized independently
- ❑ Phases communicate via interface jump relations
- ❑ Quadratic interpolation/extrapolation to interface
- ❑ Velocity-solve jump relations (1D):

$$[H] = 0$$

$$[u] = 0$$

$$[u_b] = 0$$

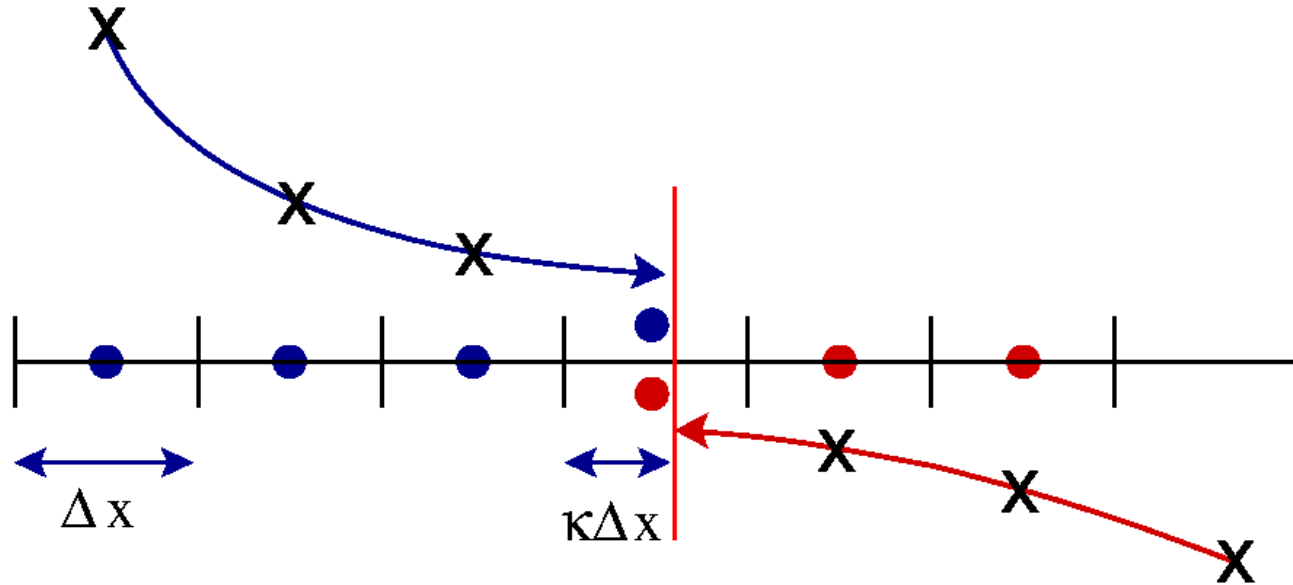
$$[\tau] = \left[\mu \frac{\partial u}{\partial x} \right] = 0$$

- ❑ System currently solved exactly (Gaussian elimination)



Multifluid Velocity Solve (cont)

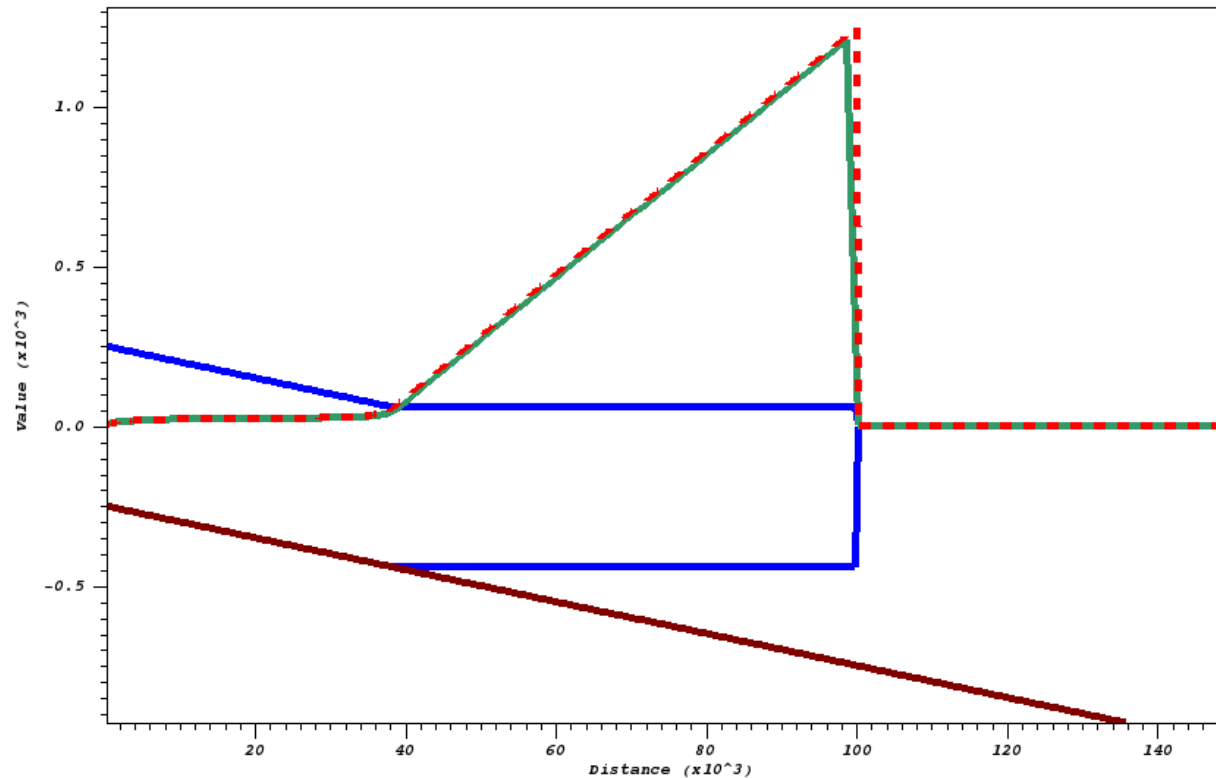
- Multifluid extrapolation to faces:



- Multivalued cell-centered value for each phase in MF cell
- Avoid “small-cell problem” ($\kappa \rightarrow 0$) by not using partial cell values in stencils
- Need quadratic extrapolant to preserve accuracy

Multifluid Velocity Solve (cont)

- ❑ Initial velocity solve
- ❑ Red dashed line: “regular” discretization $\Delta x = 195\text{m}$
- ❑ Green line, multifluid discretization, $\Delta x = 1500\text{m}$



Advection - GL advance/retreat

Two possible advection/evolution options:

1. Recompute GL every time based on finding the levelset where thickness over flotation is zero.
2. Explicitly move GL based on thickness change and basal slope

Conclusions

- ❑ Fine (sub 1-km) resolution required to get grounding lines right
- ❑ Evidence suggests that better discretizations at grounding lines may help relax resolution requirement
- ❑ Can treat GL as multifluid interfaces between 2 phases
- ❑ Multifluid velocity solver implemented
- ❑ Time-dependent evolution is next!



Acknowledgements:

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- ❑ Steph Cornford, Tony Payne at the University of Bristol
- ❑ Mark Adams (LBNL)

Extras

BISICLES Results - Ice2Sea Amundsen Sea

- ❑ Study of effects of warm-water incursion into Amundsen Sea.
- ❑ Results from Payne et al, (2012), submitted.
- ❑ Modified 1996 BEDMAP geometry (Le Brocq 2010), basal traction and damage coefficients to match Joughin 2010 velocity.
- ❑ Background SMB and basal melt rate chosen for initial equilibrium.
- ❑ SMB held fixed.
- ❑ Perturbations in the form of additional subshelf melting:
 - derived from FESOM circumpolar deep water
 - ~5 m/a in 21st Century,
 - ~25 m/a in 22nd Century.

